

Origin of Enstatite Chondrites and Implications for the Inner Planets

D. S. EBEL¹ AND C. M. O'D. ALEXANDER²

¹ Dept. of Earth & Planetary Sciences, American Museum of Natural History, NY,
NY 10024 USA (debel@amnh.org)

² Dept. of Terrestrial Magnetism, Carnegie Institution of Washington, 5241 Broad
Branch Rd., Washington DC, 20015, USA (alexande@dtm.ciw.edu)

Over 90% of EC chondrules are enstatite-rich. Some enstatite was initially Fs_5 to Fs_{30} , and was reduced after formation, probably by interaction with a vapor in which reduced matrix phases were thermodynamically stable (Lusby *et al.* 1987; Weisberg *et al.* 1994). We are interested in the nature and nebular location of such a reducing vapor, in which EC minerals (e.g. - CaS, MgS) might have been stable.

Ebel and Grossman (2000) mapped dust enrichment conditions for thermodynamic stability of FeO-rich silicates using a CI chondrite dust composition. The highly unequilibrated, anhydrous, interstellar organic- and presolar silicate-bearing cluster IDPs (C-IDPs) may be closer to the primordial dust composition than CI dust, as suggested by observed C depletions, relative to solar, in dense interstellar clouds. C-IDPs are relatively reduced, with low FeO and high C contents. Alexander (2002) noted modest enrichments of C-IDP-like dust would create conditions reducing enough to stabilize EC minerals, unless ice was also concentrated. Also, recent solar photosphere measurements suggest a 25% lower O abundance than previous estimates (Prieto *et al.* 2001).

We calculated condensation using a C-IDP-like dust composition at dust enrichments of 10, 100 and 1000 times solar at $P^{\text{tot}}=10^{-3}$ bar. Oxygen is calculated from a 75% solar baseline. The dust is H-, N-free CI, with all S as FeS, and O sufficient to make rock-forming oxides of the remaining Fe, Si, Mg, etc. At 100x enrichment, CaS is stable below 1290K, and MgS below 1180K. At 1290K, modal pyroxene (Fs_0) and olivine (Fa_0) are approximately equal. Although the system tracks $f(O_2) \sim (IW-4)$ above 1720K, $f(O_2)$ drops to (IW-8) by 1290K. Surprising! Silicate FeO decreases with decreasing T.

These results suggest that at the time the asteroids were forming, the snow line was near the inner edge of the asteroid belt, the presumed location of EC parent bodies. Bodies forming inward of the snow line would have been reduced, unless during high-T processing the dust enrichments relative to gas were modest. This has implications for the terrestrial planets' inventories of highly and moderately volatile element (and water) sensitive to reducing *versus* oxidizing conditions.

References

- Alexander C.M.O'D. (2002) *LPSC XXXIII*, Abs.#1864.
Ebel D.S. and Grossman L. (2000) *GCA* **64**, 339-366.
Lusby D., Scott, E.R.D., and Keil K. (1987) *Proc. Lunar Plan. Sci. Conf.* 17th, *J Geophys. Res. Suppl.* **92**, E679-E695.
Prieto C.A., Lambert D.L., and Asplund M. (2001) *Astrophys. J. Letters* **556**, L63-L66.
Weisberg M.K., Prinz M., and Fogel R.A. (1994) *Meteoritics* **29**, 362-373.